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LETTER

Intensifying heat extremes in China attributed to rising greenhouse gases and declining aerosols since the 2010s

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Abstract

Rising temperatures have increased the frequency, duration, and intensity of extreme temperature events over China in recent decades. The upward trends in heat extremes in China in the warm season (May-September) and their relationships with changes in aerosols and greenhouse gases are investigated using observations, reanalysis data and model results. Significant increasing trends in China are observed in daily maximum temperatures (TXx), heatwave frequency, and heatwave mean duration during 2011–2023, with increasing rates of 0.70 °C/decade, 3.77 d/decade and 0.31 d/event/decade, respectively. This study shows that 43 \pm 3% of the TXx increases in China are attributed to the rising CO₂ concentrations. Aerosol optical depth in China decreased at a rate of 0.054 per decade from 2011 to 2023 due to significant air quality improvements. The weakened aerosol cooling effect due to declining aerosols contributes $27 \pm 3\%$ to the observed TXx trend. In eastern China, where aerosol reductions were most significant, aerosol reduction even accounts for $79 \pm 10\%$ of the TXx increasing trend. The intensifying heat waves attributed to GHGs and aerosols are in accordance with the increasing extreme high temperatures. The results highlight the significant impacts of rising GHGs and decreasing aerosols on heat extreme events over China in recent years, emphasizing the need of considering both GHGs and aerosols to address the issue of intensification of heat extremes in a warming future.

1. Introduction

Due to increasing concentrations of greenhouse gases (GHGs) from human activities, warming occurs in many regions of the world, with the rise of present-day global mean surface air temperature by 1.1 °C compared to pre-industrial levels (IPCC 2021). The increases in temperature have induced unprecedented changes in the Earth system, causing rapid increases in the frequency, duration, and intensity of extreme heatwaves since the 1950s (Perkins-Kirkpatrick and

Lewis 2020). Heatwave is a specific type of extreme high temperature event, characterized by prolonged periods of excessive heat, which have been extensively applied in assessing both historical and future changes in high temperatures (e.g. Bartusek *et al* 2022, Wang *et al* 2023a). Heatwaves represent one of the most severe climate threats, with far-reaching impacts on human health, environment, and infrastructure (Singh *et al* 2021, Sharma *et al* 2022). For example, in late June 2021, an unprecedented heatwave in the Pacific Northwest region of Canada

and the United States broke the local temperature record with temperatures reaching 49.6 °C, which led to catastrophic consequences, including hundreds of deaths, marine species die-offs, reduced agricultural yields, flooding from rapid snowmelt, and a significant rise in wildfires (White *et al* 2023).

Many regions in China have observed that the frequency and intensity of heatwave events and their associated compound events in China have doubled since 1960 (Chen *et al* 2019, Luo *et al* 2020b, Wang and Yan 2021, Wang *et al* 2023b). Furthermore, persistent climate warming is anticipated to result in a substantial increase in high temperature extremes. It is projected that heatwaves would become increasingly frequent and intense under various future scenarios, according to the Coupled Model Intercomparison Project phase 6 (CMIP6) multimodel ensemble projections and machine learning methods (Zhang *et al* 2024).

Variations in GHGs and aerosols exert crucial impacts on heat extreme events. Most previous studies have primarily emphasized the role of increasing greenhouse gases (GHGs), particularly CO₂, in driving extreme temperature trends (e.g. Angélil et al 2014, Xu et al 2018). During the 20th century, the increase in anthropogenic aerosols acted to mask part of the GHGs-induced warming, especially in polluted regions like East Asia and India, leading to an underestimation of GHGs effects on extreme temperatures (Sonali and Nagesh Kumar 2016, Dileepkumar et al 2018, Seong et al 2021, Shilin et al 2024). Xu et al (2023) attributed the intensity, frequency, and duration of extreme temperatures in China during 1960-2020 in both observations and CMIP6 simulations to GHGs, aerosols, natural forcing, and land use/cover change, and found that GHGs were the main driver of extreme high temperatures changes in China, with aerosol forcing change contributing a cooling effect that partially offset the warming effect of GHGs. Zhao et al (2019) utilized the Community Earth System Model (CESM) Large Ensemble experiments to distinguish the effects of GHGs and aerosols on future global heatwaves under the Representative Concentration Pathway 8.5 scenario, showing that while more severe heatwaves came mainly from GHGs increases, aerosol reductions could have also contributed considerably over the Northern Hemisphere. Yang et al (2023) investigated the climatic impacts of aerosol reductions resulting from clean air actions in China under a carbon neutrality scenario and found that the reductions in aerosols can lead to a rise in surface temperatures in eastern China by 0.2 °C-0.5 °C in 2030 and exceeding 0.5 °C in 2060 relative to 2015. Wang et al (2023a) isolated the respective roles of

GHGs, aerosols and tropospheric ozone in driving climate change and extreme weather events under the future carbon-neutral scenario based on the CESM model simulations. They found that aerosol reductions would lead to a significant warming across the globe and a notable rise in extreme heat events, characterized by higher frequency and greater intensity, with the contributions from aerosol reductions being larger than those due to the changes in GHGs and tropospheric ozone in a carbon-neutral future.

The scientific focus of this study is to explore the trends of intensity, frequency, and duration of extreme high temperatures in China during 2011-2023 and the underlying mechanisms. The selection of this time period is motivated by the significant policy-driven reductions in emissions of aerosols and precursors in China since the 2010s, which mark a turning point from previously stable or increasing aerosol levels to a sharp decline in aerosol optical depth (AOD) across China (Benas et al 2020, de Leeuw et al 2023, Kong et al 2025, Zou et al 2025). Using observations, reanalysis data and CMIP6 multi-model simulations, this study analyzes the overall trends of extreme high temperatures in China during the warm season (May-September) from both temporal and spatial perspectives and quantifies the impact of aerosols and GHGs, respectively, on extreme temperature events.

2. Data and methods

2.1. Observations and reanalysis data

The 2-metre air temperature (T2m) data over 2000– 2023 are derived from the ERA5 dataset provided by the European Centre for Medium-Range Weather Forecasts (ECMWF), which has been utilized in many studies (e.g. Li et al 2022, Qi et al 2024). This study calculates the daily maximum temperature (Tmax) during the warm season (May–September) in China by selecting the highest value from the hourly T2m data points within each day, with a spatial resolution of 0.25° latitude/longitude. AOD data with a spatial resolution of 0.5° latitude × 0.625° longitude for the warm season during 2011-2023 are obtained from the Modern-Era Retrospective analysis for Research and Applications, Version 2 (MERRA-2) reanalysis dataset, managed by the NASA Goddard Earth Sciences (GES) Data and Information Services Center (DISC), which have been widely used in previous studies (e.g. Jia et al 2021, Jiang et al 2023, Liang et al 2024). Monthly average CO2 concentration collected from the Mauna Loa Observatory from 2011 to 2023, provided by the National Oceanic and Atmospheric Administration (NOAA) Global Monitoring Laboratory (GML), are used in this

Table 1. CMIP6-DAMIP experiments, variables and available models used in this study.

Experiment	Variable	Description	Model
hist-aer	tasmax	Daily maximum near-surface air temperature	CanESM5, BCC-CSM2-MR, FGOALS-g3, MRI-ESM2-0, MIROC6, IPSL-CM6A-LR, ACCESS-ESM1-5, ACCESS-CM2, NorESM2-LM
	od550aer	Aerosol optical thickness at 550 nm	CanESM5, MRI-ESM2-0, MIROC6, IPSL-CM6A-LR, ACCESS-ESM1-5, ACCESS-CM2, NorESM2-LM
	rtmt	Net downward radiative flux at top of atmosphere	MIROC6, IPSL-CM6A-LR, ACCESS-ESM1-5, ACCESS-CM2, NorESM2-LM
	cdnc	Cloud droplet number concentration	MIROC6, NorESM2-LM
hist-CO ₂	tasmax	Daily maximum near-surface air temperature	CanESM5, MIROC6
	rtmt	Net downward radiative flux at top of atmosphere	MIROC6

study (Keeling et al 1976). In addition, CO₂ concentration observations from the Waliguan station in China (2011–2021) (Ma et al 2020), along with daily Tmax observations from the China Meteorological Administration (CMA) during 2011–2023, are also used in this study.

2.2. DAMIP multi-model simulations

CMIP6 involves a large number of climate models, an expanded range of experimental designs, and a vast collection of simulation data (Eyring et al 2016). Previous studies have shown that CMIP6 models perform well in simulating AOD and extreme temperature changes in China, although they may slightly overestimate extreme temperatures in some regions (Li et al 2021, Zhu et al 2021, Ali et al 2022). One key activity of CMIP6 is the Detection and Attribution Model Intercomparison Project (DAMIP), which focuses on understanding contributions of various forcing factors to the observed climate change (Gillett et al 2016). DAMIP includes experiments that isolate specific drivers. The 'hist-aer' experiment examines the impact of varying anthropogenic aerosol forcing, while other forcings are fixed at their pre-industrial level. Another experiment used in this study is the 'hist-CO2' experiment, which isolates the effect of changing CO₂ concentrations from other forcings being held constant, with CO₂ used here as a representative of overall GHGs forcing due to its dominant contribution (Dong et al 2022). These targeted experiments allow researchers to better understand the individual contributions to climate variability and trends. Table 1 summarizes detailed information of the CMIP6-DAMIP experiments, variables and available models used in this study, including two historical climate experiments ('hist-aer' and 'hist-CO2')

for the period 1950–2020. To ensure consistency and maximize the sample size, we include all available models that provided the necessary outputs covering the 1950–2020 period, using the multi-model ensemble means throughout the study. This study also incorporates CO₂ concentrations data (1950–2014) from the Input4MIPs project, which provides essential input data and external forcings for CMIP6 model simulations.

The CMIP6-DAMIP multi-model simulations are employed to establish the sensitivity of heat extremes to individual external forcings. By applying linear regressions to the CMIP6-DAMIP 'hist-aer' and 'hist-CO₂' experiments over 1950–2020, we obtain the responses of heat extreme indices to changes in AOD and CO₂ concentration, respectively. The long-term period starting from year 1950 helps reduce internal variability and ensures more robust estimates. These derived sensitivities are then combined with observed AOD and CO2 trends during 2011-2023 to estimate their contributions to recent changes in heat extremes. This approach effectively links modelbased responses with real-world observations and follows established detection and attribution practices (Bindoff et al 2014, Jones et al 2016, Seong et al 2022). In addition, we analyze cloud and radiation variables from the 'hist-aer' and 'hist-CO2' experiments to investigate the physical impacts of aerosol reductions and GHGs increases since the 2010s. This helps to further understand the mechanisms behind the observed changes in heat extremes.

2.3. Extreme high temperature indices

To measure the impacts of anthropogenic forcings on extreme high temperature events, two types of extreme indices, TXx and heatwave-related indices,

are derived from Tmax during the warm season in China and analyzed in this study. TXx is defined as the highest Tmax of the year for each grid, as a key metric to represent the intensity of extreme high temperature events (e.g. Wang et al 2017, Iles et al 2024). Following many previous studies on extreme heat events (e.g. Zhang et al 2020, Wang et al 2023a), a heatwave event in this study is defined as a period of at least 3 consecutive days with Tmax exceeding the 90th percentile threshold based on the 2000-2023 climatological average at each grid point, in order to establish a reliable status with a sufficiently long dataset. Heatwaves are characterized by the total number of heatwave days and the average duration of heatwave events during the warm season, serving as indicators of the frequency and duration of extreme high temperature events.

3. Results

Figure 1 presents the linear trends of heat extreme indices across China during the warm season from 2011 to 2023, based on ERA5 reanalysis data. TXx shows a general increasing pattern across most of China (figure 1(a)), at a rate up to 0.40 °C per year in parts of the North China Plain and Sichuan Basin, which is validated by CMA station observations (figure S1). Regions with statistically significant warming trends are concentrated primarily in central, southwestern, southern and eastern China. During 2011-2023, more than 98% of days experiencing TXx in China are during heatwaves. The trends in heatwave indices show substantial increases in the frequency and duration of heatwaves in many regions of China (figures 1(b) and (c)), although some areas, particularly in western and northeastern China, exhibit negative trends, possibly due to local climate variability such as changes in soil moisture, cloud cover, precipitation or circulation patterns (e.g. East Asian summer monsoon), which could suppress surface warming or heatwave persistence (Whan et al 2015, Wei et al 2023, Ye et al 2025). Moreover, the relatively short analysis period and lack of observational data in western China may also affect the regional trends. The nationally averaged extreme temperature indices in China during 2011-2023 also present significant increases, with TXx rising at 0.70 °C/decade, heatwave frequency increasing at 3.77 d/decade, and heatwave mean duration increasing at 0.31 d/event/decade (figures 1(d)–(f)). Overall, these trends illustrate a general pattern of increasing maximum temperatures, more frequent heatwave days, and longer heatwave durations in China during the warm season from 2011 to 2023, underscoring the warming climate and the increasing impact of extreme heat events in China over the recent decade.

Over the past few decades, the rising concentrations of GHGs, particularly CO₂, have been a key driver of global warming, which has directly

influenced the increasing extreme temperature events (Kweku et al 2018, Xu et al 2023). Recent studies have emphasized the complex role of aerosols in the global climate system, revealing that reduced emissions of aerosols and precursors have diminished their historical cooling effect, thereby increasing surface air temperatures and decreasing relative humidity, with this relative warming effect potentially surpassing that of GHGs in certain scenarios (Zhao et al 2019, Wang et al 2023a). Assessing the impacts of changes in GHGs and aerosols on heat extreme trends in China requires an understanding of the relationship between these forcings and heat extreme indices, which can be derived from the CMIP6-DAMIP experiments. In this study, AOD and CO₂ concentrations are employed as metrics to measure variations in aerosols and GHGs, respectively.

Figure 2 illustrates the relationships between extreme temperature indices and the two influencing factors in China during the warm season, based on the CMIP6-DAMIP 'hist-aer' and 'hist-CO₂' experiments from year 1950 to present-day. In the historical period, AOD constantly increased in China during 1950-2009 and then slightly decreased in 2010-2020 (figures 2(a)-(c)), while CO_2 concentrations continued increasing since 1950 (figures 2(d)–(f)). Increased AOD, primarily associated with anthropogenic aerosols, cools the atmosphere by reducing sunlight reaching the surface and modifying cloud properties (Boucher et al 2013, Bellouin et al 2020). The cooling effect from rising AOD reduces extreme temperatures and heatwave frequency and duration, as evidenced by their negative correlations. This relationship remains robust during the period of declining AOD after 2010 (figures S2). Based on the linear regressions, a unit increase in AOD would reduce TXx by 3.5 ± 0.5 °C, heatwave days by 10.55 ± 1.76 d yr⁻¹, and heatwave mean duration by 2.15 \pm 0.43 d/event in China during the warm season, as derived from the CMIP6-DAMIP 'hist-aer' experiment. In contrast, increases in CO2 concentrations are linked to the rising extreme temperature indices, with an increase of 0.012 \pm 0.001 °C in TXx, 0.041 \pm 0.008 d yr⁻¹ in heatwave days and 0.008 \pm 0.002 d/event in heatwave mean duration in China for each 1 ppm rise in CO₂, as simulated in the CMIP6-DAMIP 'hist-CO₂' experiment.

Figure 3 shows the time series of warm season CO₂ concentrations and AOD in China during 2011–2023 from observations and reanalysis data. Both CO₂ concentrations from NOAA GML and at Waliguan station show a monotonic growth, with trends of 24.67 ppm/decade and 24.04 ppm/decade, respectively. According to the sensitivities of heat extreme indices to CO₂ concentrations derived from the 'hist-CO₂' experiment (figures 2(d)–(f)), the CO₂ increase translates into a TXx trend of about 0.3 \pm 0.02 °C/decade, heatwave frequency of 1.01 \pm 0.20 d/year and heatwave mean duration

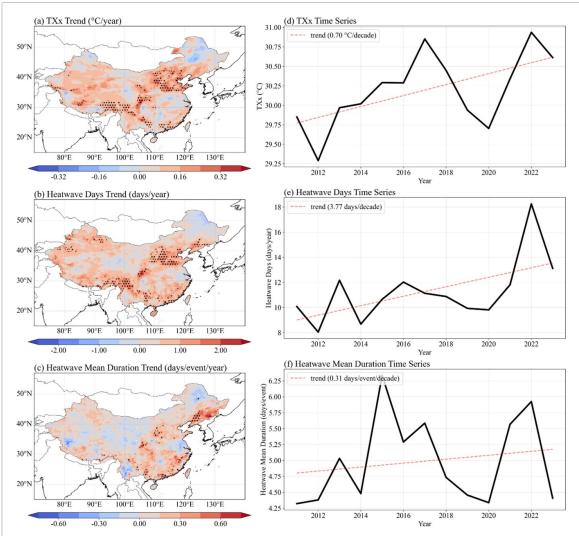


Figure 1. Linear trends in heat extreme indices in China from ERA5 reanalysis data. Spatial distribution of linear trends of (a) TXx ($^{\circ}$ C, the highest Tmax of the year for each grid), (b) heatwave days (days/year), and (c) heatwave mean duration (days/event) in the warm season (May–September) in China during 2011–2023. Time series (solid lines) of (d) TXx ($^{\circ}$ C), (e) heatwave days (days/year), and (f) heatwave mean duration (days/event) in the warm season in China from 2011 to 2023, with dashed lines representing the linear fitted trends. The shaded areas in a, b and c denote statistically significant at the 95% confidence level using the Student's t test.

of 0.2 \pm 0.05 d/event. It suggests that the increase in CO₂ concentration contributed 43 \pm 3% to the 0.7 °C/decade of the observed increasing trend of TXx in China during 2011–2023, as well as 27 \pm 5% and 65 \pm 16% of the increases in heatwave frequency and heat wave mean duration, respectively, due to their positive correlation with CO₂ concentrations.

The AOD in China shows a significant decline from 2011 to 2023, at a rate of -0.054 per decade. The reduction is largely attributed to the implementation of China's clean air policies, which have effectively curbed aerosol pollution. Based on the relationships between heat extreme indices and AOD derived from the 'hist-aer' experiment (figures 2(a)-(c)), the decline in AOD contributed approximately $27\pm3\%$ to the observed TXx trend during 2011-2023, as well as $15\pm3\%$ and $37\pm7\%$ of the increasing heatwave frequency and heatwave mean duration, respectively. These findings suggest that while CO_2 concentrations remain the dominant driver of heat extreme

temperature increases, the reduction in aerosols from air quality improvements has also played a significant role in the intensifying heat extremes.

Aerosols affect surface temperature by directly scattering and absorbing sunlight, which reduces solar radiation reaching the surface (aerosolradiation interaction), and indirectly by changing cloud properties, leading to more, smaller droplets that brighten clouds and increase cloud cover, thus cooling the surface (aerosol-cloud interaction) (Twomey 1974, Albrecht 1989, Luo et al 2020a, Huusko et al 2022). GHGs like CO2 mainly warm the surface by trapping outgoing longwave radiation, enhancing the greenhouse effect (Ramanathan et al 1985, Chen and Dong 2019, Raghuraman et al 2023). Therefore, decreases in aerosols reduce their cooling effect, while rising CO₂ concentration increases warming, together influencing extreme heat events. The reduction in aerosols due to clean air actions leads to decreases in AOD and cloud

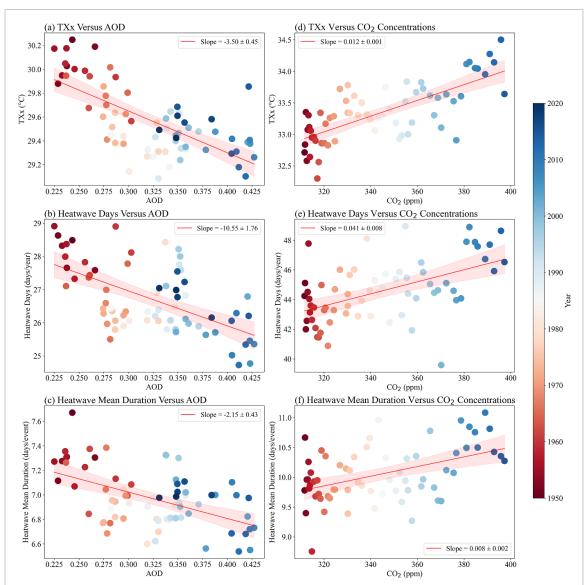


Figure 2. Relationships between heat extreme indices and AOD (unitless) or CO_2 concentrations (ppm) in China. Scatter plot of multi-model ensemble mean (a) TXx (°C), (b) heatwave days (days/year), and (c) heatwave mean duration (days/event) versus AOD in the warm season during 1950–2020 from the CMIP6-DAMIP 'hist-aer' experiment. Scatter plot of multi-model ensemble mean (d) TXx (°C), (e) heatwave days (days/year), and (f) heatwave mean duration (days/event) from the CMIP6-DAMIP 'hist- CO_2 ' experiment versus CO_2 concentrations from input4mips data in the warm season during 1950–2014. Lines represent the slope of linear fitting with standard errors, and the 95% confidence intervals are shown as shaded areas. The p-values of all fitted relationships are less than 0.01.

droplet number concentration (CDNC), which through aerosol-radiation and aerosol-cloud interactions cause a significant increase in the top-of-atmosphere (TOA) net downward radiative flux (figures 4(a)–(c)). It indicates that more radiation reaches the Earth, enhancing warming and intensifying extreme heat events in China. At the same time, the sustained increase in CO_2 concentration primarily enhances the absorption of longwave radiation emitted from the surface, resulting in an increase in TOA net downward radiative flux and thereby generating a greenhouse effect (figure 4(d)). This process is mainly driven by radiative forcing, but the associated warming may induce cloud feedbacks

that moderately modulate the net radiative effect. Together, aerosol reductions and rising CO₂ concentrations act through distinct radiative mechanisms to jointly drive increases in the frequency, intensity, and duration of extreme high-temperature events.

Regional attributions in eastern China (east of 105°E) are also examined here, considering that this region not only experienced the high levels of aerosol loading but also exhibited the rapid decline in AOD across China since 2011 (figure S3). From 2011 to 2023, AOD in eastern China decreased at a rate of -0.13 per decade, while TXx, heatwave frequency, and heatwave mean duration increased

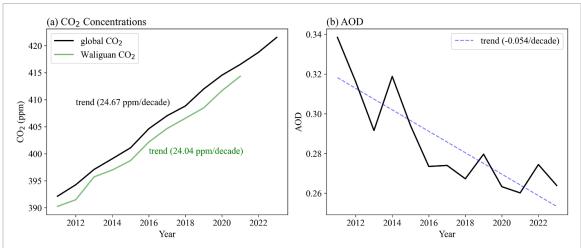


Figure 3. CO_2 concentrations (ppm) and AOD (unitless) trends in China. (a) Time series of global CO_2 concentrations from NOAA GML during 2011–2023 (black line) and CO_2 concentrations at the Waliguan station during 2011–2021 (green line). (b) Time series of regionally averaged warm season mean AOD in China from MERRA-2 reanalysis during 2011–2023, with dashed lines representing the linearly fitted trends.

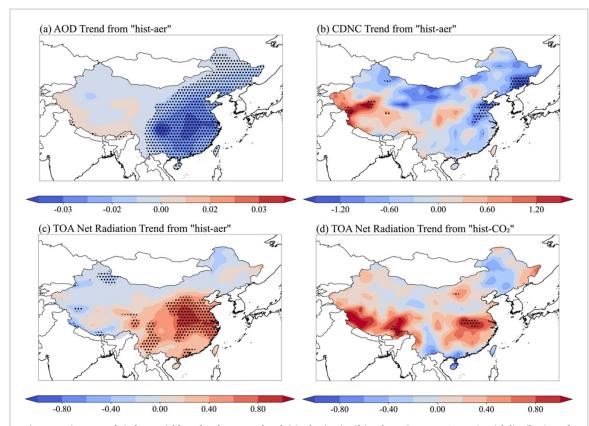


Figure 4. Linear trends in key variables related to aerosol and CO_2 forcing in China from CMIP6-DAMIP. Spatial distribution of trends in multi-model ensemble mean (a) AOD (/year), (b) CDNC (cm⁻³ yr⁻¹), (c) TOA net downward radiative flux (W m⁻² yr⁻¹) from the CMIP6-DAMIP 'hist-aer' experiment, and (d) TOA net downward radiative flux (W m⁻² yr⁻¹) from the CMIP6-DAMIP 'hist-CO₂' experiment in China during 2010–2020, with shaded areas representing statistically significant at the 95% confidence level using the Student's t test.

at a rate of 0.58 °C/decade, 3.41 d/decade and 0.64 d/event/decade, respectively (figure 5). The reductions in AOD contributed 79 \pm 10% to the observed trend in TXx, 40 \pm 7% to the trend in

heatwave frequency, and 44 \pm 9% to the trend in heatwave mean duration. In contrast, the increase in CO₂ concentrations contributed approximately 52 \pm 3% to the TXx trend, 30 \pm 6% to the heatwave

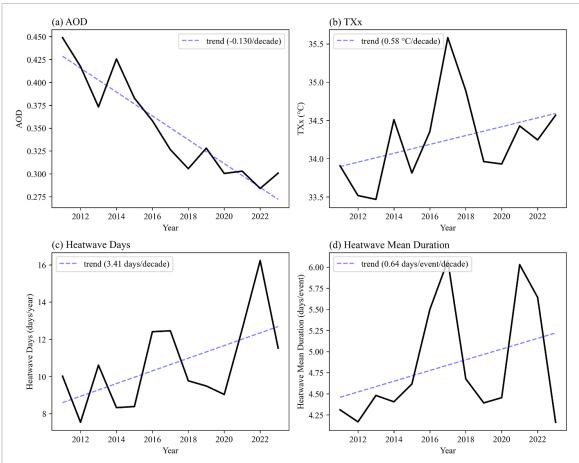


Figure 5. Linear trends in MERRA-2 AOD and heat extreme indices from ERA5 reanalysis data over eastern China (east of 105°E). Time series of regional averaged (a) mean AOD (unitless), (b) TXx (°C), (c) heatwave days (days/year), and (d) heatwave mean duration (days/event) over eastern China during the warm season from 2011 to 2023, with dashed lines representing the linearly fitted trends.

frequency trend, and $31\pm8\%$ to the heatwave mean duration trend. The results show a critical role of aerosol reductions played in driving the observed changes in extreme heat events in eastern China, contributing to higher intensity, more frequency, and longer duration of heatwaves.

4. Conclusions and discussions

This study reveals a significant intensification of heat extreme events across China during the warm season (May–September) from 2011 to 2023, characterized by rising maximum temperatures as well as more frequent and endured heatwave events. The core motivation of our study is to quantify the impact of the substantial decline in aerosol emissions in China since 2010s on the intensification of heat extremes, and to assess the relative contributions of aerosol reductions and rising GHGs to this recent trend, a dimension not explicitly addressed in previous literatures. The observed increasing trends in heat extreme indices, particularly TXx, heatwave frequency, and heatwave

mean duration, are consistent with the broader patterns of global warming, driven primarily by the increasing concentrations of GHGs and decreasing aerosols. Our analysis shows that the rise in CO₂ level has exerted a profound influence on these heat extreme indices, contributing 43 \pm 3%, 27 \pm 5% and $65 \pm 16\%$ to the observed increasing trends in TXx, heatwave frequency and heatwave mean duration, respectively. The reduction in aerosols contributed approximately 27 \pm 3% to the observed warming in TXx, $15 \pm 3\%$ to the increasing heatwave frequency, and 37 \pm 7% to the prolonged heatwave duration from 2011 to 2023. Aerosol reductions decrease AOD and CDNC, enhancing TOA net downward radiative flux via aerosol-radiation and aerosol-cloud interactions, while rising CO₂ increases longwave radiation absorption, strengthening the greenhouse effect. In eastern China, the impacts of aerosol reductions were even more pronounced, with the reduction of AOD accounting for 79 \pm 10% of the warming trend in TXx, $40 \pm 7\%$ of the increase in heatwave frequency, and $44 \pm 9\%$ of the rise in heatwave mean duration,

which surpass the contributions of CO₂, especially in terms of maximum temperatures and the heatwave frequency. It suggests that air quality improvements, while being beneficial to public health, may have unintended consequences for climate extremes, particularly in a warming world.

In conclusion, while efforts of reducing CO₂ emissions are critical to addressing the climate change, strategies to manage aerosol concentrations also hold significant implications for mitigating heat extreme events. As China continues to combat air pollution, the trade-off between improving air quality and mitigating climate extremes must be carefully considered, especially in regions experiencing substantial aerosol reductions. Further research is needed to refine the understanding of these interactions and to explore effective strategies for minimizing the impacts of both aerosols and GHGs on heat extreme events in the future.

Data availability statement

The data that support the findings of this study are openly available at the following URL/DOI: https://doi.org/10.5281/zenodo.15172975 (Yang 2025).

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